



Hydrogeology of surficial, unconsolidated quaternary aquifers, Maynard Creek catchment, Bridger Range, Montana  
by John F Whittingham

A thesis submitted in partial fulfillment of the requirements for the degree of Master of Science in Earth Sciences  
Montana State University  
© Copyright by John F Whittingham (1996)

Abstract:

The groundwater regime within the Maynard Creek catchment is poorly defined. Aquifers may exist in fractured bedrock and surficial, unconsolidated Quaternary deposits. Ski area and residential development within the catchment has increased water demands. This investigation addresses the following questions: 1) Do the surficial, unconsolidated Quaternary deposits function as aquifers?; 2) Can surficial aquifer discharge be distinguished from bedrock aquifer discharge?; and 3) Can the surficial aquifers be characterized? The techniques used in this investigation include: 1) geologic mapping; 2) stream hydrograph and flow duration curve analyses; 3) groundwater discharge monitoring; 4) hydrogeochemical analyses; 5) material sampling; and 6) seismic refraction.

Geologic mapping was completed to delineate potential aquifers. Groundwater discharge features, including 62 springs, are strongly clustered within surficial deposits. Most springs and seeps, and all perennial streamflow exist at lower elevations where surficial deposits are thickest. Stream hydrograph analyses indicate that discharge is sustained by direct surface water runoff and that groundwater discharge is limited.

Springs within surficial deposits display rapidly increased yields after intense precipitation and snowmelt runoff. Three perennial springs exist and each discharges from surficial deposits. These springs yield between 3 lps (0.1 cfs) and 28 lps (1.0 cfs).

Spring water temperatures decrease with increased elevation by 1.0° (1.8°F) per 100 m (328 ft). Groundwater discharges from surficial deposits have low winter temperatures (2 to 4°C, 36 to 39°F) that steadily increase through the summer. The average temperature variability of the most persistent surficial springs is 1.9°C (3.4°F). Water temperatures of bedrock springs are higher (5.0°C, 41°F), with an average variability of 0.2°C (0.4°F). Water temperatures of discharge from surficial deposits are rapidly depressed 0.1 to 1.0°C (0.2 to 1.8°F) by meteorological and snowmelt runoff events. The most productive perennial spring's annual water temperature cycle mimics the annual mean daily air temperature cycle with a lag time of 60 days.

Specific electrical conductance of all 62 spring waters was 0.16 to 0.50 (mS). This indicates low ionic concentrations of total dissolved solids and short residence times for stored groundwater. Springs flowing from surficial deposits discharge Ca (calcium) and HCO<sub>3</sub> (bicarbonate) enriched waters. Well waters derived from bedrock are Na (sodium) and HCO<sub>3</sub> (bicarbonate) enriched.

Material sampling and seismic refraction indicate that rock glacier till, undifferentiated till, outwash, and buried channels are the primary components of surficial aquifers. Porosity, storativity, and hydraulic conductivity were approximated for Quaternary materials. Rock glacier till, outwash, and buried channels yield rapid and intermittent groundwater discharge. Undifferentiated tills sustain perennial stream baseflows. The surficial Quaternary aquifers have development potential, but excessive winter withdrawals will reduce perennial stream baseflows.



HYDROGEOLOGY OF SURFICIAL, UNCONSOLIDATED  
QUATERNARY AQUIFERS, MAYNARD CREEK  
CATCHMENT, BRIDGER RANGE, MONTANA

by

John F. Whittingham

A thesis submitted in partial fulfillment  
of the requirements for the degree

of

Master of Science

in

Earth Sciences

MONTANA STATE UNIVERSITY  
Bozeman, Montana

May 1996

© COPYRIGHT

by

John Frederic Whittingham

1996

All Rights Reserved

N378  
W6187

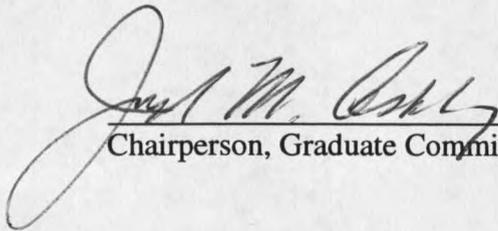
APPROVAL

of a thesis submitted by

John Frederic Whittingham

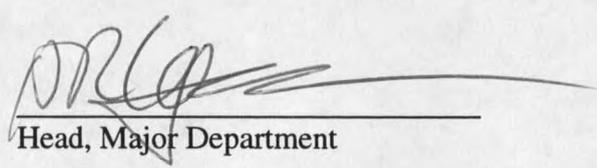
This thesis has been read by each member of the thesis committee and has been found to be satisfactory regarding content, English usage, format, citations, bibliographic style, and consistency, and is ready for submission to the College of Graduate Studies.

21 Jan 1996  
Date

  
Chairperson, Graduate Committee

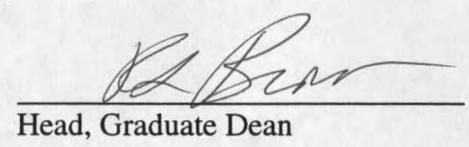
Approved for the Major Department

21 Jan 1996  
Date

  
Head, Major Department

Approved for the College of Graduate Studies

3/2/96  
Date

  
Head, Graduate Dean

## STATEMENT OF PERMISSION TO USE

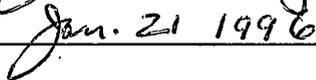
In presenting this thesis (paper) in partial fulfillment of the requirements for a master's degree at Montana State University, I agree that the Library shall make it available to borrowers under the rules of the Library.

If I have indicated my intention to copyright this thesis (paper) by including a copyright notice page, copying is allowable only for scholarly purposes, consistent with "fair use" as described in the U.S. Copyright Law. Requests for permission for extended quotations from or reproduction of this thesis (paper) in whole or in parts may be granted only by the copyright holder.

Signature



Date



## VITA

John Frederic Whittingham was born to David and Helen Whittingham in Williamsburg, Virginia on November 6, 1959. He attended elementary through high school grades in Colorado Springs, Colorado, and graduated from Thomas B. Doherty High School in 1977. Mr. Whittingham attended the University of Colorado in Boulder from 1977 to 1979 and participated in general studies and Earth Sciences. Between 1979 and 1981, Mr. Whittingham completed Emergency Medical Technician certification at Pikes Peak Community College in Colorado Springs, and pursued Ski Area Management and avalanche studies at Colorado Mountain College, in Leadville and Breckenridge, Colorado. He was a member of the El Paso County, Colorado, Search and Rescue between 1979 and 1981. In 1981, Mr. Whittingham enrolled at the University of Montana in Missoula. In 1986, he completed a B.A. in Geology with a professional emphasis in Hydrogeology, and a B.A. in Geography with a corollary field in Watershed Management. Mr. Whittingham worked as a teacher, builder, and geologist from 1986 to 1990. In 1990, Mr. Whittingham initiated graduate studies in Earth Sciences at Montana State University, in Bozeman. He completed a M.S. in Earth Sciences in 1996. He also completed the Center for High Elevation Studies program requirements at Montana State University. In 1992, Mr. Whittingham worked as a hydrologist technician for the United States National Forest Service, Beaverhead National Forest, in Dillon, Montana. In 1993, Mr. Whittingham began permanent employment as a hydrologist with the Bureau of Land Management, Dillon Resource Area, in Dillon, Montana. Mr. Whittingham is currently employed (1996) by the Bureau of Land Management in Dillon, where he administers the Soil, Water, and Air, and Hazardous Materials Programs.

## TABLE OF CONTENTS

	Page
LIST OF TABLES .....	vii
LIST OF FIGURES .....	viii
LIST OF PLATES .....	xi
ABSTRACT .....	xii
INTRODUCTION .....	1
Geographic Setting .....	3
Structural Geology .....	3
Stratigraphy .....	6
Hydrogeologic Setting .....	6
Hydrologic Setting .....	9
The Problem .....	10
Purpose .....	10
METHODS .....	11
Mapping .....	11
Aerial Photographs and Existing Maps .....	11
Geologic Mapping .....	12
Geomorphologic Mapping .....	13
Hydrogeologic Mapping .....	14
Aquifer Material Characterization .....	15
Surficial Materials and Potential Aquifers .....	15
Existing Research .....	16
Seismic Refraction .....	16
Site Examination .....	17
Hydrology .....	18
Existing Data .....	18
Stream Discharge Measurements .....	18
Water Chemistry .....	19
Existing Data .....	19
Field and Laboratory Methods .....	19
Spring Water Monitoring .....	20
Spring Inventory and Monitoring .....	20
Spring Discharge .....	21
Specific Electrical Conductance .....	22
Spring Water Temperatures .....	22

TABLE OF CONTENTS - Continued

	Page
RESULTS .....	23
Mapping .....	23
Geologic Mapping .....	23
Geomorphologic Mapping .....	28
Hydrogeologic Mapping .....	35
Aquifer Material Characterization .....	37
Surficial Materials and Potential Aquifers .....	37
Existing Research .....	37
Seismic Refraction .....	37
Site Examination .....	45
Hydrology .....	49
Existing Data .....	49
Stream Discharge Measurements .....	54
Water Chemistry .....	55
Existing Data .....	55
New Data .....	57
Spring Water Monitoring .....	57
Spring Inventory and Monitoring .....	57
Spring Discharge .....	58
Specific Electrical Conductance .....	58
Spring Water Temperatures .....	60
INTERPRETATIONS .....	62
Surficial, Unconsolidated Quaternary Aquifers .....	62
Interpretation of Mapping Results .....	62
Hydrograph Analyses .....	67
Spring Water Temperatures .....	70
Surficial Quaternary Aquifers and Bedrock Aquifers .....	77
Field Interpretations .....	77
Chemical Interpretations .....	78
Spring Water Monitoring Interpretations .....	82
Characterization of Surficial, Unconsolidated Quaternary Aquifers .....	85
Quaternary Deposits that Do Not Function as Aquifers .....	85
Glacial Deposits as Aquifers .....	87
Bridger Aquifer .....	91
Deer Park Aquifer .....	95
Groundwater Development .....	103
CONCLUSION .....	108
APPENDIX 1 .....	115
REFERENCES CITED .....	123

## LIST OF TABLES

Table	Page
1. Published Estimated Values of Porosity, Storativity, and Hydraulic Conductivity for Unconsolidated Materials .....	38
2. Seismic Velocities and Calculated Depths of Materials in the Maynard Catchment .....	42
3. Published Seismic Velocities of Geologic Materials .....	43
4. The Range and Average Seismic Velocities of Specific Geologic Materials in the Maynard Creek Catchment. ....	46
5. Calculated Monthly Discharge of Maynard Creek .....	51
6. Chemical Analyses of Waters Collected from Wells Located in the Maynard Creek Catchment .....	55
7. Chemical Analyses of Waters Collected from Springs and Creeks from Within or Near the Maynard Creek Catchment. ....	56
8. Chemical Analyses of Waters Collected from Perennial Springs in the Maynard Creek Catchment. ....	57
9. Ionic Balance of Chemical Analyses Used During this Investigation .....	80
10. Inferred Values of Porosity, Storativity, and Hydraulic Conductivity for Surficial, Unconsolidated Quaternary Materials Inspected in the Maynard Creek Catchment .....	103

## LIST OF FIGURES

Figure	Page
1. The Bridger Range and Maynard Creek Catchment . . . . .	2
2. Topographic Map of the Maynard Creek Catchment . . . . .	4
3. Generalized Stratigraphic Section . . . . .	7
4. The Bridger Groundwater Discharge Investigation Area . . . . .	39
5. The Deer Park Groundwater Discharge Investigation Area . . . . .	39
6. Time-Distance Diagram for Seismic Line 8: Depiction of a Two-Layer Stratigraphy . . . . .	40
7. Time-Distance Diagram for Seismic Line 17: Depiction of a Three-Layer Stratigraphy . . . . .	40
8. Time-Distance Diagram for Seismic Line 14: Depiction of a Two-Layer Stratigraphy and a Low-Velocity Discontinuity . . . . .	41
9. Time-Distance Diagram for Seismic Line 19: Depiction of a Two-Layer Stratigraphy and a Low-Velocity Discontinuity . . . . .	41
10. Frequency and Distribution of Seismic Velocities of the Materials Located in the Bridger and Deer Park Groundwater Discharge Areas . . . . .	43
11. Examples of Texture of Two Separate Undifferentiated Tills . . . . .	48
12. Outwash (Left) and Buried Deposits (Right) at a Springhead . . . . .	48
13. Stream Hydrograph for Maynard Creek . . . . .	49
14. Location of Sites, Facilities, and Measurement Points . . . . .	50
15. Snow-Water Equivalent of Snowpack: Maynard Creek SNOTEL Station, Water Years 1990 and 1991 Compared to 30-Year average (1961-1990) . . . . .	52

LIST OF FIGURES - Continued

Figure	Page
16. Snow-Water Equivalent of Snowpack: Bridger Bowl SNOTEL Station, Water Years 1990 and 1991 Compared to 30-Year average (1961-1990) . . . . .	53
17. Mean Daily Air Temperature Measured at the United States Weather Bureau 12 NE Station, Maynard Creek Catchment . . . . .	53
18. Stream Discharge (l/s) for the North and South Forks of Maynard Creek . . . . .	54
19. The Discharge Class Distribution of Springs at Three Different Times of Observation . . . . .	59
20. Deer Park Spring Hydrograph (Spring 23). . . . .	59
21. Deer Park Spring Temperature Variability, April 20, 1991 Through June 7, 1992 . . . . .	60
22. Temperature Variability of Springs Located at Similar Elevations . . . . .	61
23. Three Different Spring Water Temperature Regimes . . . . .	61
24. The Frequency of Springs in Each Morphologic Category . . . . .	64
25. Flow Duration Curve for Maynard Spring Stream Discharge . . . . .	68
26. Correlation of All Spring Water Temperatures and Elevation . . . . .	71
27. Correlation of Spring Water Temperature and Elevation . . . . .	72
28. The Mean Daily Air Temperature at the Bozeman 12 NE Station and Spring Water Temperature of the Deer Park Spring (Spring 23) . . . . .	74
29. Snow-Water Equivalent at the Maynard Creek SNOTEL Station and Discharge from the Deer Park Spring (Spring 23) . . . . .	75
30. Temperature and Discharge of the Deer Park Spring (Spring 23) . . . . .	76

LIST OF FIGURES - Continued

Figure	Page
31. Piper Diagram of the Chemical Composition of Select Waters . . . . .	81
32. Temperature Regimes of Springs Sustained by Bedrock Aquifer Storage (Springs 51 and 53) and Surficial Quaternary Aquifer Storage (Springs 23, 48, and 49) . . . . .	83
33. Soil Water Discharge Model . . . . .	86
34. Rock Glacier Aquifer Configuration . . . . .	88
35. Glacial Outwash and Undifferentiated Till Aquifer Model . . . . .	89
36. Undifferentiated Till Aquifer Model . . . . .	91
37. Schematic Interpretation of Seismic Line 14 . . . . .	93
38. Schematic Interpretation of Seismic Line 17 . . . . .	94
39. The Extent and Approximate Thickness of the Deer Park Aquifer . . . . .	96
40. Cross-Section of the Deer Park Aquifer Including Approximate Values of Storativity (S) in Percent, Hydraulic Conductivity (K) in m/d, and Baseflow Discharge (Qb) in lps . . . . .	96
41. Response Delay in Spring Discharge to Recharge in Glacial Soils . . . . .	98
42. Schematic Interpretation of Seismic Line 19 . . . . .	99
43. Schematic Interpretation of Seismic Line 3 . . . . .	100
44. Schematic Interpretation of Seismic Line 4 . . . . .	102

LIST OF PLATES

Plate	Page
1. Geologic Map . . . . .	insert, back page
2. Surficial Geologic Map . . . . .	insert, back page
3. Surface Hydrology Map . . . . .	insert, back page

## ABSTRACT

The groundwater regime within the Maynard Creek catchment is poorly defined. Aquifers may exist in fractured bedrock and surficial, unconsolidated Quaternary deposits. Ski area and residential development within the catchment has increased water demands. This investigation addresses the following questions: 1) Do the surficial, unconsolidated Quaternary deposits function as aquifers?; 2) Can surficial aquifer discharge be distinguished from bedrock aquifer discharge?; and 3) Can the surficial aquifers be characterized? The techniques used in this investigation include: 1) geologic mapping; 2) stream hydrograph and flow duration curve analyses; 3) groundwater discharge monitoring; 4) hydrogeochemical analyses; 5) material sampling; and 6) seismic refraction.

Geologic mapping was completed to delineate potential aquifers. Groundwater discharge features, including 62 springs, are strongly clustered within surficial deposits. Most springs and seeps, and all perennial streamflow exist at lower elevations where surficial deposits are thickest. Stream hydrograph analyses indicate that discharge is sustained by direct surface water runoff and that groundwater discharge is limited.

Springs within surficial deposits display rapidly increased yields after intense precipitation and snowmelt runoff. Three perennial springs exist and each discharges from surficial deposits. These springs yield between 3 lps (0.1 cfs) and 28 lps (1.0 cfs).

Spring water temperatures decrease with increased elevation by 1.0°C (1.8°F) per 100 m (328 ft). Groundwater discharges from surficial deposits have low winter temperatures (2 to 4°C, 36 to 39°F) that steadily increase through the summer. The average temperature variability of the most persistent surficial springs is 1.9°C (3.4°F). Water temperatures of bedrock springs are higher (5.0°C, 41°F), with an average variability of 0.2°C (0.4°F). Water temperatures of discharge from surficial deposits are rapidly depressed 0.1 to 1.0°C (0.2 to 1.8°F) by meteorological and snowmelt runoff events. The most productive perennial spring's annual water temperature cycle mimics the annual mean daily air temperature cycle with a lag time of 60 days.

Specific electrical conductance of all 62 spring waters was 0.16 to 0.50 (mS). This indicates low ionic concentrations of total dissolved solids and short residence times for stored groundwater. Springs flowing from surficial deposits discharge Ca (calcium) and HCO<sub>3</sub> (bicarbonate) enriched waters. Well waters derived from bedrock are Na (sodium) and HCO<sub>3</sub> (bicarbonate) enriched.

Material sampling and seismic refraction indicate that rock glacier till, undifferentiated till, outwash, and buried channels are the primary components of surficial aquifers. Porosity, storativity, and hydraulic conductivity were approximated for Quaternary materials. Rock glacier till, outwash, and buried channels yield rapid and intermittent groundwater discharge. Undifferentiated tills sustain perennial stream baseflows. The surficial Quaternary aquifers have development potential, but excessive winter withdrawals will reduce perennial stream baseflows.

## INTRODUCTION

Groundwater storage and availability in steep mountainous terrain is highly variable and often limited by geologic constraints. In southwestern Montana, groundwater development in Upper Cretaceous rocks often produces low yields and poor water quality, while several Mississippian through Lower Cretaceous formations provide excellent groundwater quantity and quality (Miller, 1974, Moore, 1984, Ferreira and others, 1986). Overlying Quaternary deposits also may function as aquifers, but often they are clayey and do not transmit useful amounts of water, or lack sufficient storage volume to fulfill even modest human needs. In most cases, inadequate data exists for the hydrogeologic regime of mountain aquifer systems to be fully understood.

In the Maynard Creek catchment, on the eastern flank of the Bridger Range (Figure 1), steeply dipping Paleozoic through Upper Cretaceous bedrock is overlain by an undetermined thickness of unconsolidated Quaternary materials. Summer observations reveal an abundance of springs and seeps, and vigorous stream flow. During winter, few springs flow and stream discharge is minimal. The nature of groundwater occurrence in the catchment is unknown.

Resort development plans and increasing recreational activities in and around the Maynard Creek catchment have resulted in demands for additional consumptive and nonconsumptive water uses. Surface water availability from Maynard Creek is severely limited. Surface water is overappropriated and typical summer discharge is less than existing claimed water rights (Morrison-Maierle, Inc., 1978). Groundwater may be a viable resource to supplement the current and future water demands.

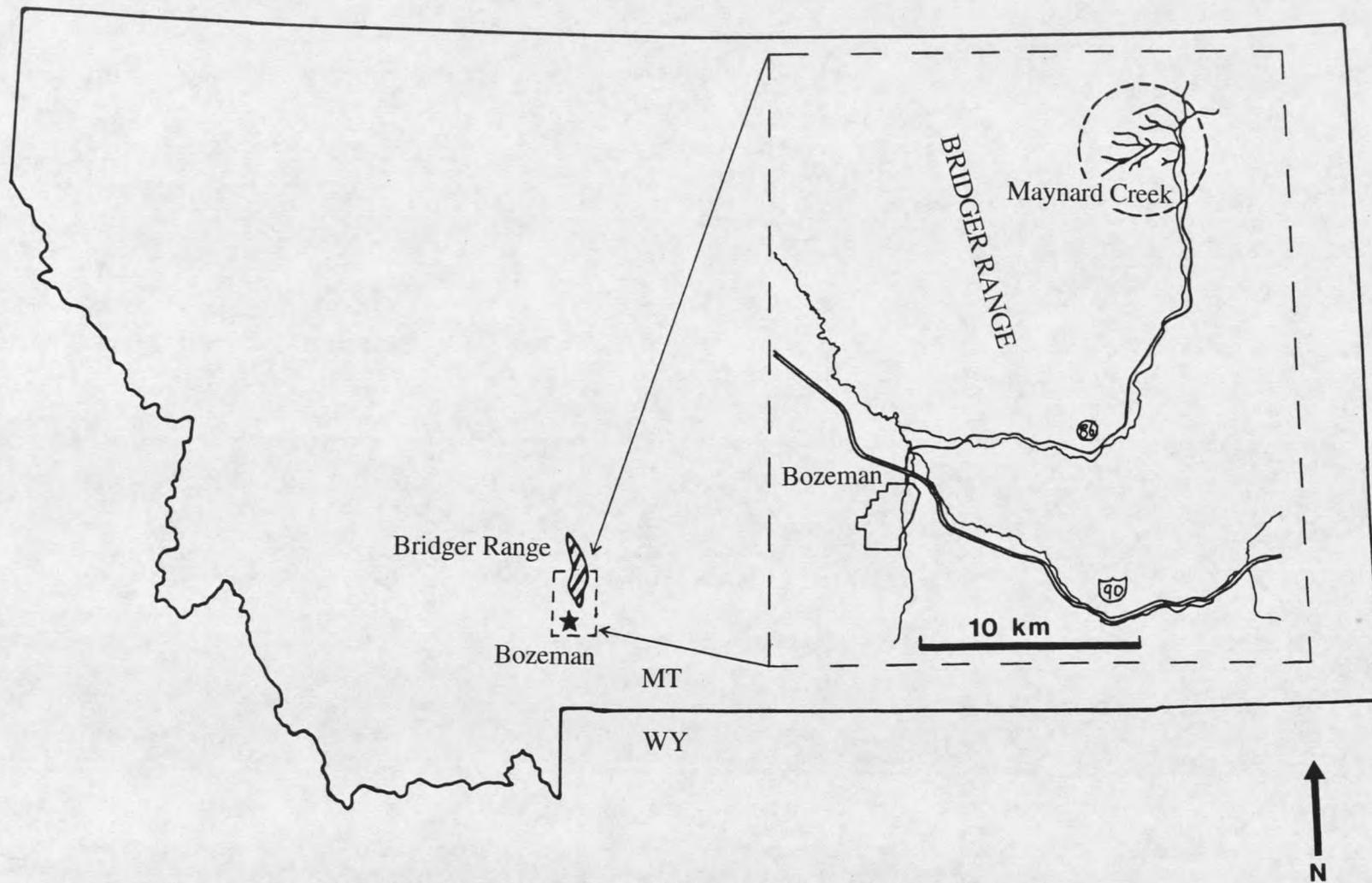


Figure 1. Location of the Bridger Range and the Maynard Creek catchment.

The problem of groundwater development in the catchment is that potential bedrock and surficial Quaternary aquifers have not been identified, and the capacity for water production from these materials is unknown. The purpose of this investigation is to determine the significance of surficial Quaternary aquifers. The objectives are to identify surficial, unconsolidated Quaternary aquifers; to distinguish surficial, unconsolidated Quaternary aquifer waters from bedrock aquifer waters; and to characterize the components of the surficial, unconsolidated Quaternary aquifer system in the Maynard Creek catchment.

### Geographic Setting

The Maynard Creek catchment is on the eastern flank of the central Bridger Range, approximately 26 km (16 mi) northeast of Bozeman, Montana (Figure 2). Maynard Creek flows into Bridger Creek, which is a main tributary of the East Gallatin River. The catchment is east facing and has 883 m (2897 ft) of local relief, rising from 1774 m (5820 ft) at Bridger Creek to 2657 m (8717 ft) on the crest of the Bridger Range. A total of 5.6 km<sup>2</sup> (2.2 mi<sup>2</sup>) are contained within the catchment boundaries (Figure 2).

### Structural Geology

The structural geology of the Bridger Range has been described by McMannis (1955) and Lageson (1989). The Range is a small north-trending uplift which is approximately 40 km (25 mi) in length. It is located about 80 km (50 mi) east of the Boulder batholith, and constitutes a part of the Northern Rocky Mountain Front. The Bridger Range is unique in that it overlaps the boundaries of four major tectonic provinces. These tectonic provinces are: 1) the Middle Proterozoic Belt Basin; 2) the Sevier fold and thrust belt; 3) the Laramide foreland province of basement involved deformation; and 4) the Basin and Range province







































































































































































































































































